

Geodynamic setting of fluxoturbidites in West Carpathian Upper Jurassic and Lower Cretaceous sedimentary basins

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Abstract. Eustatic and hydrodynamic regime on Tethyan shelves and microcontinents changed during Late Jurassic and Early Cretaceous. These changes are well recognizable by study of the allodapic fluxoturbidite intercalations in pelagic basinal infillings.

Tithonian (lithohorizon Ti-6) Barmstein fluxoturbidites with platform carbonate clasts are restricted to S of the Alpine - Carpathian lithospheric block. No biohermal rims supplying allodapic beds have been developed around contemporaneous Jasenina Basin of northernmore Fatic Zone. Nozdovica Breccia (lithohorizon Be-7) accumulated along active faults: its composition changed from place to place.

Valanginian (Oravice, Va-4) siliciclastites inserted in hemipelagites (equivalent to East Alpine Rossfeld Beds) originated during eustatic sea level drop by river transport bringing terrigenous clastic material from SE. Chromian spinel grains were derived from eroded ultrabasic rocks of elevated segments of fundament.

The maximum development of fluxoturbidites has been attained during Hauterivian (Ha-3). The Strážovce Turbidite consists of bioherm rim debris, transported by channel system cutting fault slope.

Barremian (Ba-2) and Aptian (Ap-2) fluxoturbidites were associated with prograding Urganian carbonate platforms in southern Tatic area. Other allodapes were derived from eroded elevated substrate.

Albian siliciclastites accumulated in turbidite fan rim arising during extensive paleogeographic re-building of Central Western Carpathians. The quartzose admixture was derived from erode fundament and, partially, it was also produced by syngenetic volcanism.

Key words: calciturbidites, siliciclastites, pelagic carbonate basins, Upper Jurassic, Lower Cretaceous, Western Carpathians.

The intercalation of allodapic sediments in pelagic Upper Jurassic and Lower Cretaceous sequences was associated with an eustatic drop of the sea-level and with the effects of tectonic sea bottom denivelisation.

Although the Neo-Cimmerian collision zone running meridionally south of Asian Cimmerides terminates in Crimea and Dobrogea, traces (both compressive and tensional) of Neo-Cimmerian deformations are recognizable in the Carpathians (mainly in their southernmost zones: MICHALÍK 1990) and in the Eastern Alps: (TOLLMANN 1987), as well.

The deformation of outer Tethyan shelves changed the eustatic and hydrodynamic regime of the basins. During the Berriasian, global transgression (VAIL et al. 1977) and the successive change of current regime led to paleogeographic reorganisation of the broader area of the Mediterranean Tethys (MICHALÍK et al. 1991, SOTÁK 1989). Reef growth on the Outer Carpathian border finished, the extent of carbonate platforms diminished substantially: their products have been preserved only rarely, mostly in fragments in allodapic sediments. Microplankton remnants became an important constituent of pelagic sediments (VAŠÍČEK et al. 1994). Nannocone biomicrites sedimented in aerated eupelagic environment, while radiolarites and silicites were deposited below the CCD level. Turbidite beds and other allodapic members are inserted in Upper Jurassic and Lower Cretaceous pelagic carbonate successions of Central and Outer Western Carpathian basinal infillings. A wide range of facies types in the sections of Zliechov, Bobrovec, Vysoká, Belá, Havran and Humenné units of the Križna Nappe, in the Butkov development of the Manin Unit (all in Central Carpathians), as well as in the Rača Unit of the Magura Basin (Outer Carpathians) allows to follow lateral transitions from distal through proximal calciturbidite deposits to deep sea fans, slope lithofacies and infillings of submarine channels.

Oxfordian to Berriasian allodapes were formed of the detritus of platform carbonates (REHÁKOVÁ 1995) transported down the slopes of elevated fundament blocks (MICHALÍK 1994). Fluxoturbidites (Barmstein Limestone) are restricted to the basin bottom in southern zones of the Alpine - Carpathian lithospheric block (MIŠÍK & SÝKORA 1982). Fluxoturbidite beds are accompanied by intercalations of contourite character.

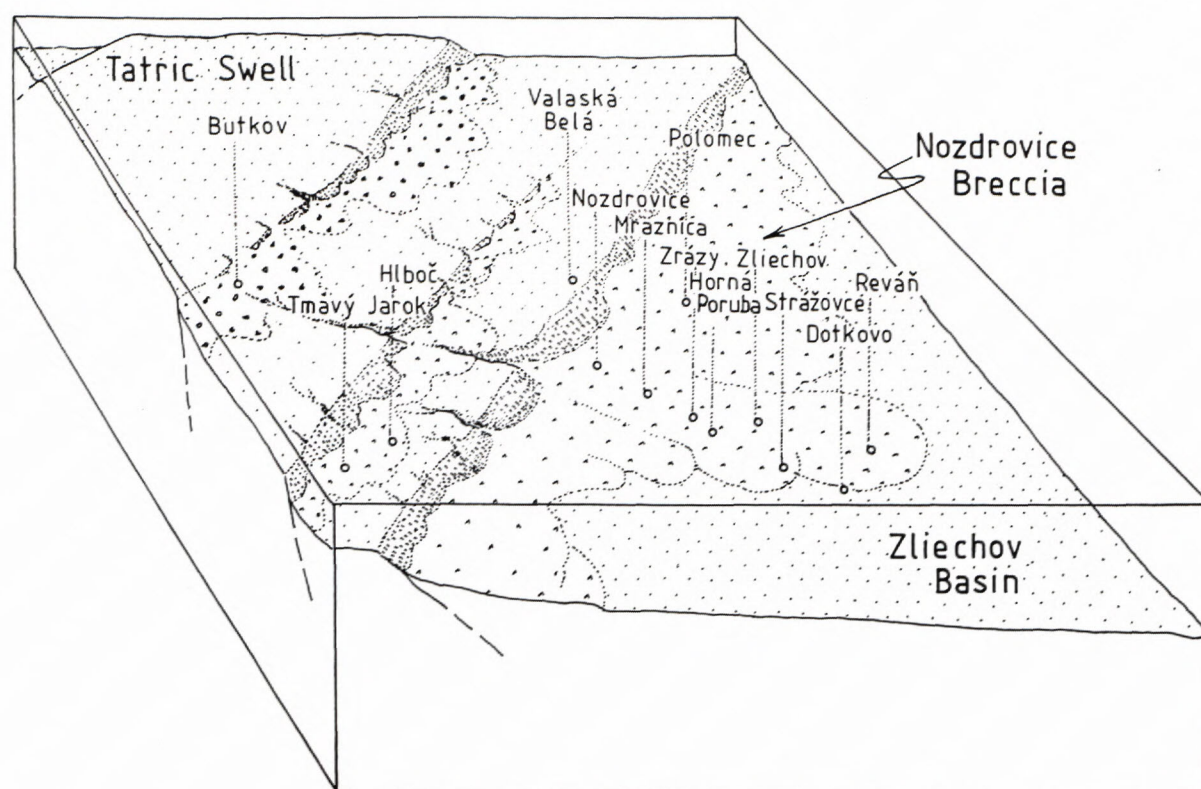


Fig. 1. Model of origin of the Nozdovice Breccia Beds in NW part of the Fatric Zliechov Basin during Berriasian time.

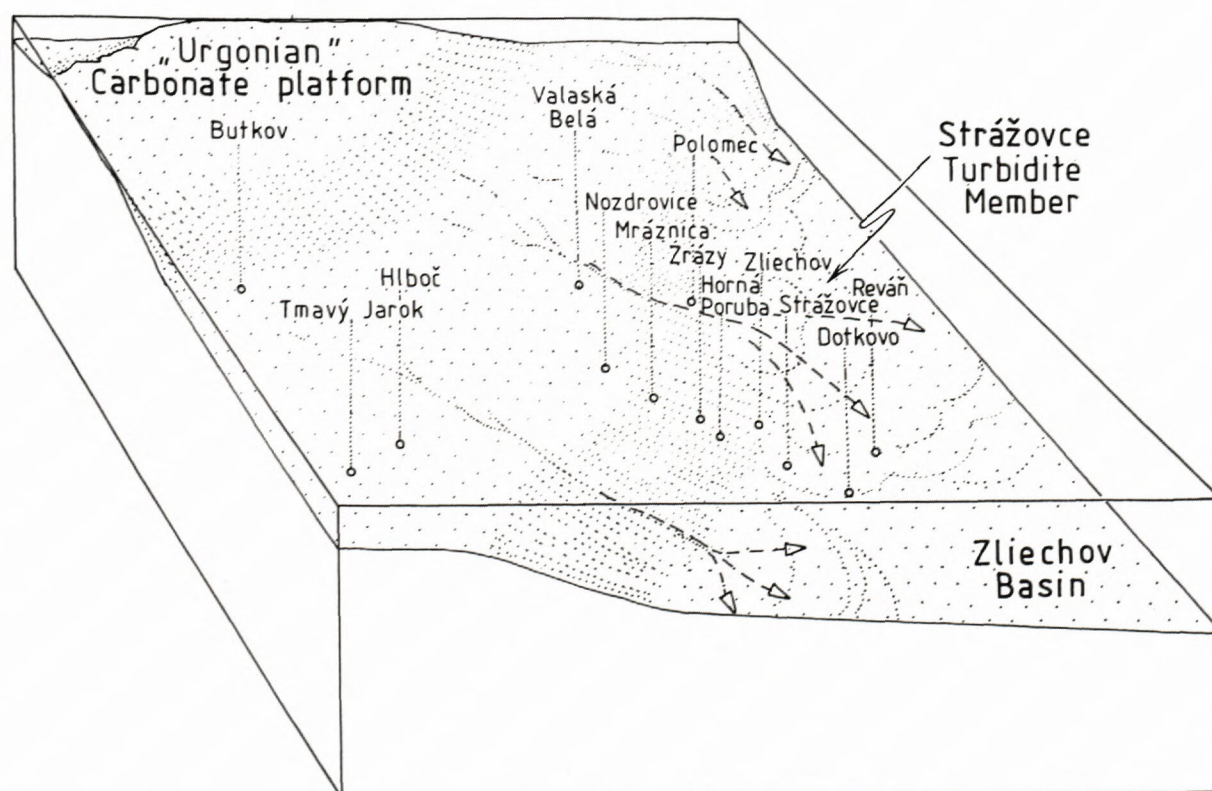


Fig. 2. Model of origin of the Strážovce Turbidite Member in NW part of the Fatric Zliechov Basin during Hauterivian time.

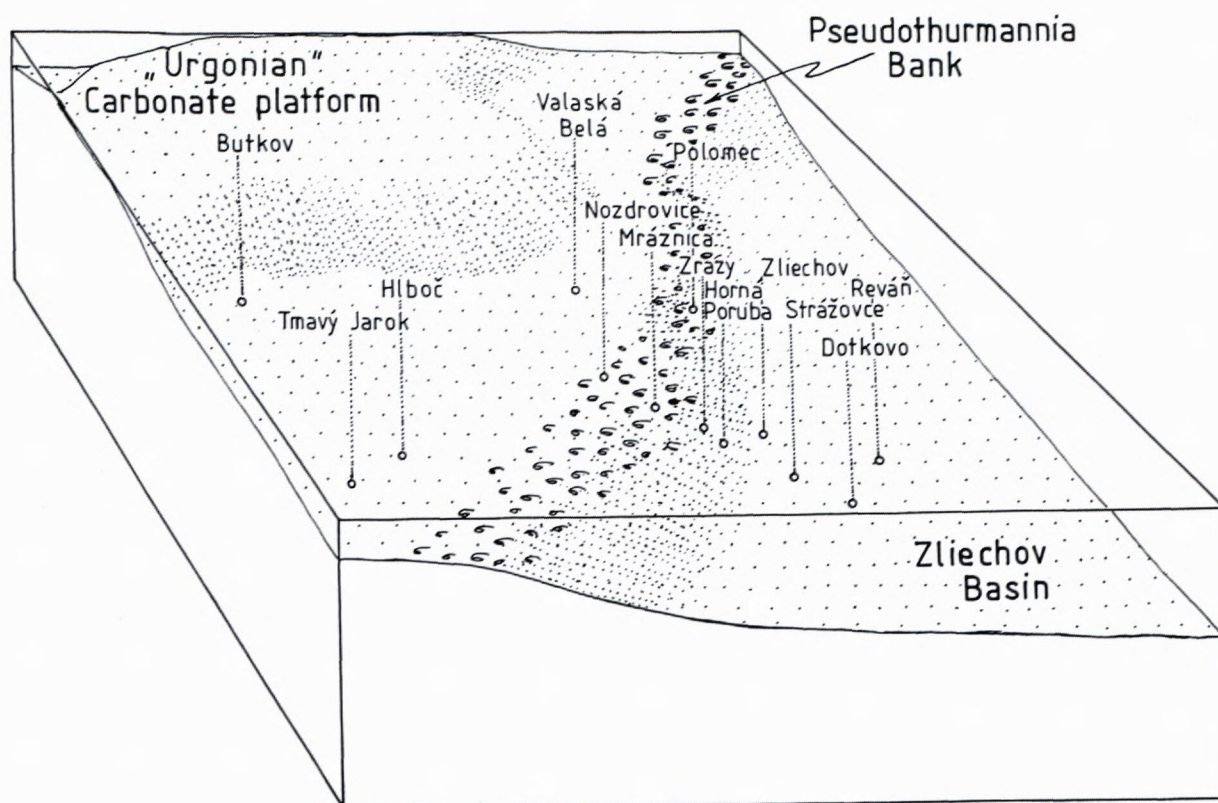


Fig. 3. Model of origin of the *Pseudothurmannia* Bed in NW part of the Fatric Zliechov Basin during latest Hauterivian time.

On the other hand, the contemporaneous filling of the Zliechov Basin (Jasenina Formation) in the Fatricum does not contain any allodapic beds. Apparently, this basin was not rimmed by bioherms at that time. Bodies of the Nozdovica Breccia composed of basemental pelagic Upper Jurassic/ Berriasian limestone clasts accumulated along active faults (Fig.1; MICHALÍK & REHÁKOVÁ, 1995). Proximal developments (Butkov, Hlboč, Nozdovica localities) have more variegated composition of clasts, while the breccia along the axis of the assumed submarine fan (Mrázňica, Zrážy, Poruba, Zliechov, Strážovce, Reván localities) is composed of Upper Tithonian *Crassicollaria* and Lower Berriasian *Calpionella* limestone clasts.

Upper Berriasian and Valanginian pelagic sequence contains allodapic intercalations as well. However, the maximum development of these facies was attained during Hauterivian. Calciturbidite intercalations are preserved in Magura Basin infilling in Outer Carpathians (VAŠÍČEK & REHÁKOVÁ 1994). Lower Cretaceous resediments in Pieniny Klippen Zone units occur rather exceptionally (AUBRECHT 1994).

BORZA et al.(1980) defined a Hauterivian organodetrital limestone complex inserted in the Mrázňica Formation as the Strážovce Turbidite Member. It consists of limy detritus of the bioherm rim, transported into basin by a

channel system cutting the fault slope along its border (Fig.2.). During the topmost Hauterivian shallowing and Tatic carbonate platform progradation, this border was the place of accumulation of fossiliferous *Pseudothurmannia* Beds (Fig.3).

The Strážovce Member is well exposed by the Polomec quarries near Lietavská Lúčka (BORZA et al., 1984). Probably the thickest body of allodapes (Murán Limestone Formation) forming a big deep sea fan lens transported southwards into basinal carbonates of the Koscieliska and Muránska Lúčka Formations builds the Belianske Tatry Mountains Ridge (MICHALÍK & SOTÁK 1990).

On the other hand, silicic turbidite beds contain material derived from different sources. Cimmerian deformed Inner Carpathian zones on the SE of the Alpine - Carpathian microcontinent could be regarded as a possible source area. Several silicic turbidite beds are inserted in the Valanginian hemipelagic Fatric sequence in the Juráňova Valley near Oravice (JABLONSKÝ in AUBRECHT et al., 1992). They are genetically equivalent to the clastic Rossfeld Beds in the Eastern Alps (FAUPL & TOLLMANN 1979). Their origin was connected with an eustatic drop of sea level, with the activity of rivers bringing terrigenous clastic material into carbonate basins at that time, as well as with tectonic activity of individual segments of the

fundament. The presence of chromium spinel grains derived from eroded ultrabasic rocks is remarkable (Mišík et al., 1980, Mišík & SÝKORA, 1981, FAUPL & WAGREICH 1992, WAGREICH et al., 1995).

The rising proportion of radiolarians and sponge spicules in Valanginian rocks could culminate in local concentrations of these remnants. Although some of them were caused by environmental fluctuations, others could be accumulated by distal turbidity or countour currents. Similar thin parallel bedded calcisiltites with radiolarian bands were described from the Lombardian Basin of the Southern Alps (WEISSERT, 1981; BERSEZIO, 1993).

Barremian and Aptian fluxoturbidites were frequently associated with prograding biogenic sedimentation on Urgonian carbonate platforms (MICHALÍK & SOTÁK 1990; MICHALÍK 1994) in southern Tatric area. Other allodapic deposits were derived from eroded elevated parts of the substrate (JABLONSKÝ et al. 1993).

Allodapic beds of siliciclastic character occur more frequently in Aptian and Albian sequences. JABLONSKÝ & MARSCHALCO (1992) suggested extensive paleogeographic re-building of the area of Central Western Carpathians during the middle Albian. It was connected with basin subsidence and with the uplift of northwestern marginal zones. This process led to the accumulation of a turbidite fan rim. The quartzose admixture was derived from the eroded fundament and, partially, it was also produced by syngenetic volcanism.

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